



**ETH**

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# Revised snow scheme in the ECMWF land surface model: Offline validation and impacts on EC-EARTH

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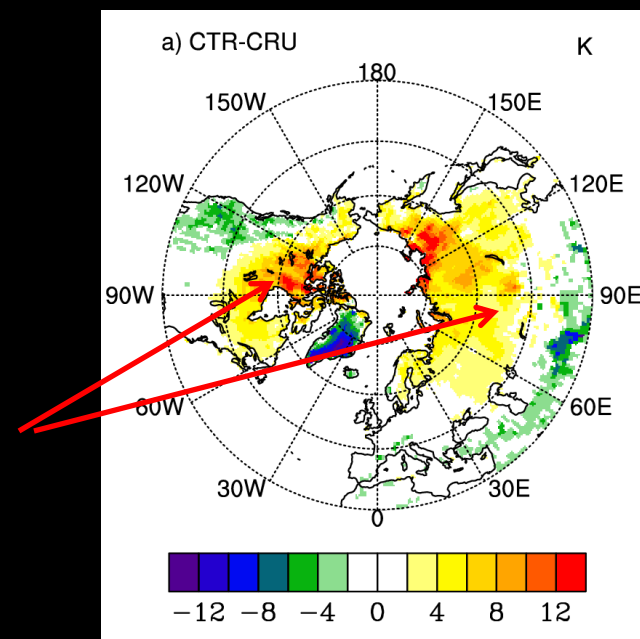
Thanks to the EC-EARTH consortium members, Simona Stefanescu and Camiel Severijns. Snowmip2 data providers, Richard Essery, Nick Rutter, Martin Hirschi and Crystal Schaaf.

Why is snow important ? weather forecast and Climate modeling

- **Surface characteristics** : changes in the surface albedo, roughness etc....
- **Processes/feedbacks** : Snow albedo feedback (Chess et al 1991, Science, etc...)  
Impact N. Hemisphere circulation (Gong et al, 2007 G.P.C, etc...)  
Indian summer Monsoon (Robock et al 2003 JGR, etc...)
- **Weather forecast** : errors in the “physics” are attenuated by data assimilation;
- **Climate modeling** : assimilation is not an option;

• **Local site simulations (offline) and climate (coupled) runs pointed some deficiencies in the ECMWF land surface scheme (HTESSEL) .**

• **Is the snow scheme responsible for the warm bias over snow-covered regions ?**



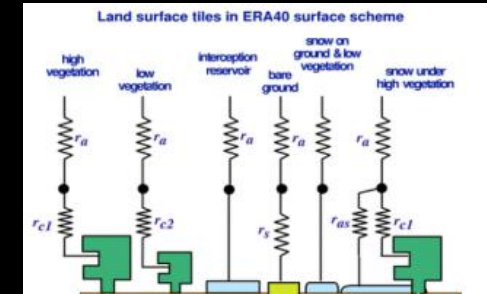
Winter 2 meter temperature biases in the ECMWF model in climate mode (EC-EARTH)

# HTESSEL and NEW snow scheme

## HTESSEL (Hydrology - Tiled ECMWF Scheme of Surface Exchanges over Land)

Balsamo et al. 2009 J.H

- Up to 5 surface tiles (bare ground, low and high vegetation, interception,) + ocean and sea ice (binary land sea mask);
- **Two tiles for snow:**
  - **exposed snow;**
  - **shaded snow (under high vegetation).**



**CTR**

**OPER** (Dutra et al. 2010 J.H)

<b>Liquid water</b>	1-Dry snow 2-Rainfall bypass the snowpack	1-Diagnosed from snow mass and temperature 2-Interception of rainfall 3-Changed snow energy and mass balance
<b>Snow Density</b>	1-Empirical exponential increase 2-Snowfall density constant=100 kg.m <sup>-3</sup>	1-Physically based (Anderson, 1976) 2- Snowfall density : function of temperature and wind speed
<b>Snow Albedo</b>	1- Exponential(melting) / Linear decay 2- Reset to max (0.85) snowfall > 1 mm hr <sup>-1</sup> 3- Shaded: constant = 0.15	1- Exponential (when liquid water is present) / Linear decay; 2- Continuous reset to max depending on the amount of snowfall (10 mm to full reset) 3-Shaded : vegetation type dependent (Moody et al. 2007)
<b>Snow fraction</b>	1-Function of snow mass 2- Threshold (SC=1) : 15 mm	1-Function of snow mass and density 2- Threshold (SC=1): 10 cm (10 mm -> 40 mm)

# Simulations setup and validation

Increasing computational cost

## 1) Offline - local sites (HTESSEL)

- SnowMip2 (5 sites x 2 seasons - Rutter et al. 2009 JGR)

Process validation

## 2) Offline - global (HTESSEL)

- GSWP2 (1°x1° 1986-1995)
- ERA-Interim (0.7° 1989-present)

Global validation  
(only land surface)

## 3) Operational runs (IFS: TL255)

- Extended data assimilation
- Hindcasts : 1 year climate runs

Short range forecast

## 4) Climate runs (EC-EARTH TL159)

- Atmosphere only
  - SSTs and sea ice from observations
  - 30 years run (1979-2008) 3 member ensemble
- Coupled ocean (not shown)

Feedbacks

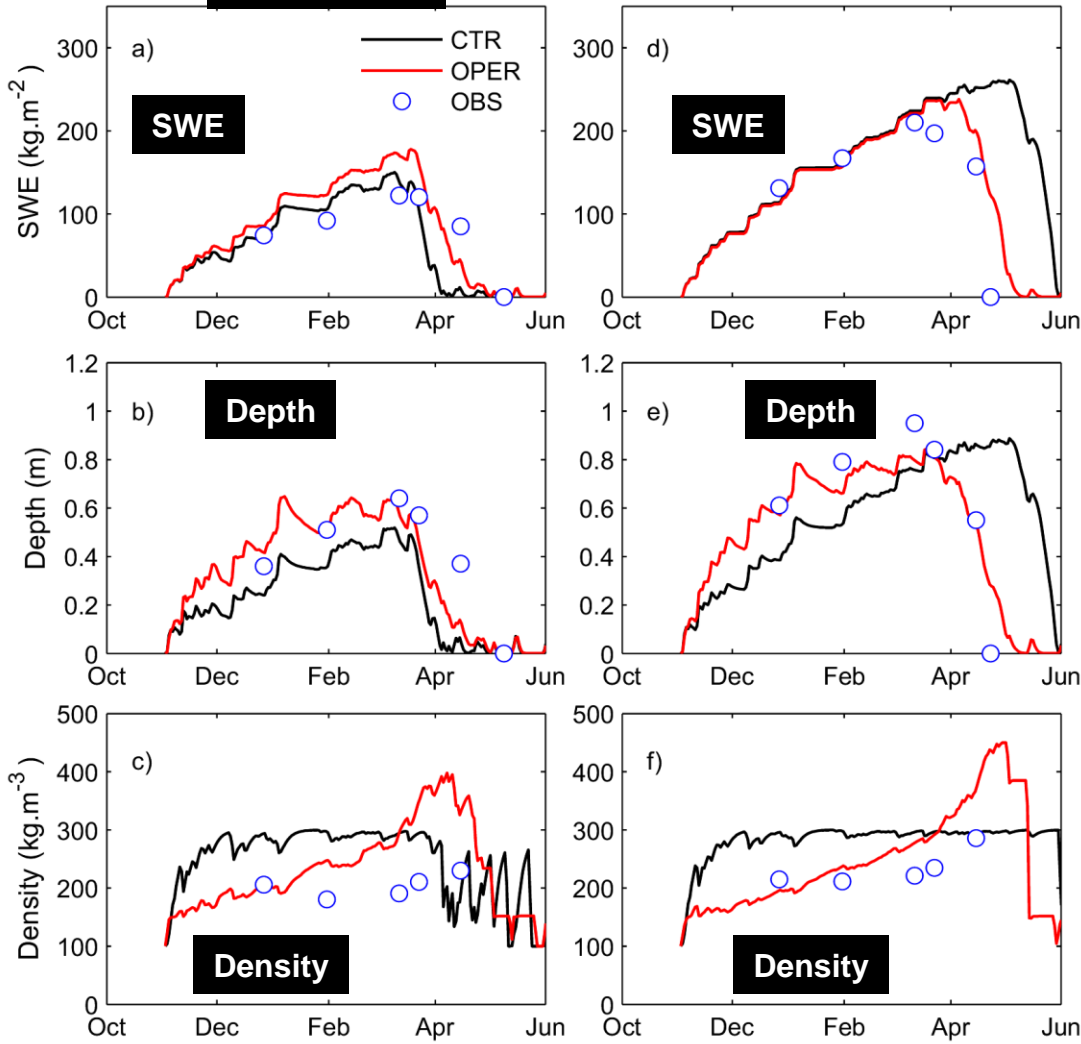
# Snowmip2 (site) simulations

## Fraser (Colorado USA)

Forest site

Fraser 2003-04

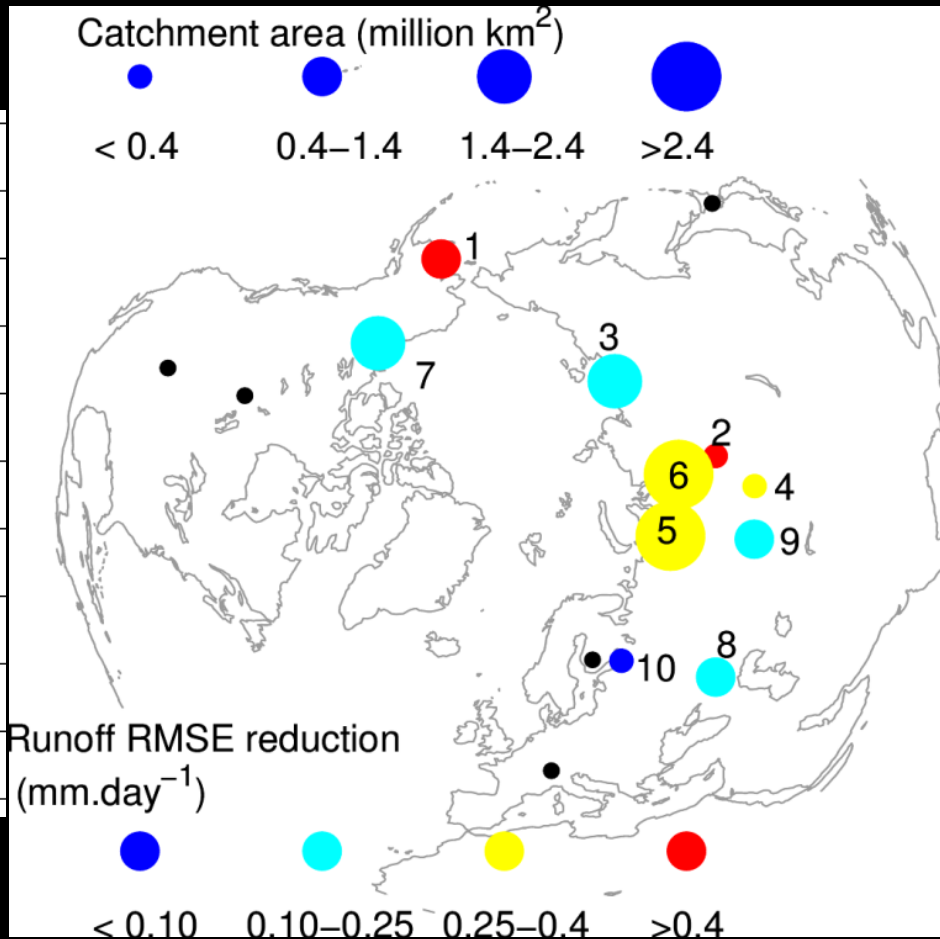
Open site



- Early melting in forest sites
  - 21\* days – CTR, 13\* days OPER
- Late melting in open sites
  - 10\* days – CTR, 2\* days OPER
- \* Averaged SnowMip2 sites (5 sites x 2 seasons)
- Snow density:
  - Exponential increase in CTR
  - Closer to observations in NEW

# Basin scale validation

1-Yukon
2-Podka. <sup>a</sup>
3-Lena
4-Tom
5-Ob
6-Yenisei
7-Mackenzie
8-Volga
9-Irtish
10-Neva



- Reduced snow density
- Increased snow insulation
- Less soil cooling;
- Less soil freezing;
- Increased water infiltration;
- Reduction of surface runoff
- Increased bottom drainage (lagged in time);
- Improved timing of late Spring/Summer runoff
- Increased soil water storage

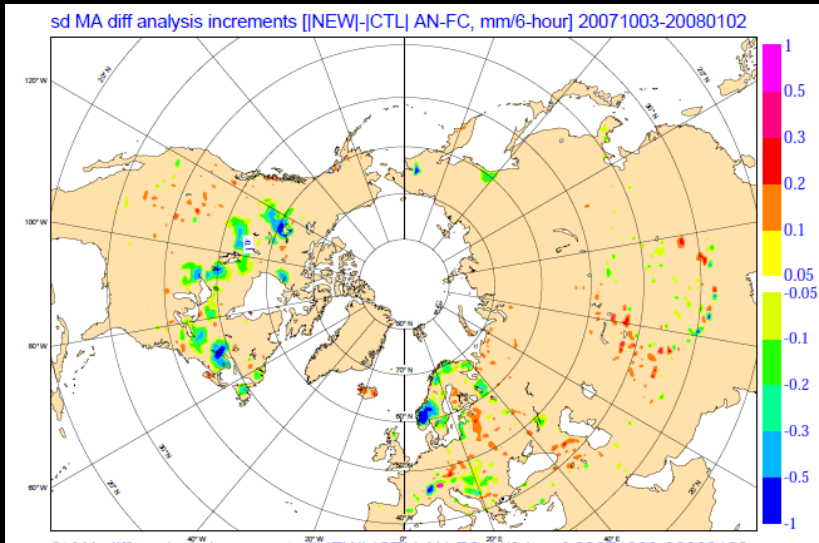
Average RMSE of runoff mm day<sup>-1</sup>  
**CTR:0.75**  
**OPER:0.51 ( reduction of 32%)**

BSWB (runoff data): Hirschi et al. 2006, J.H

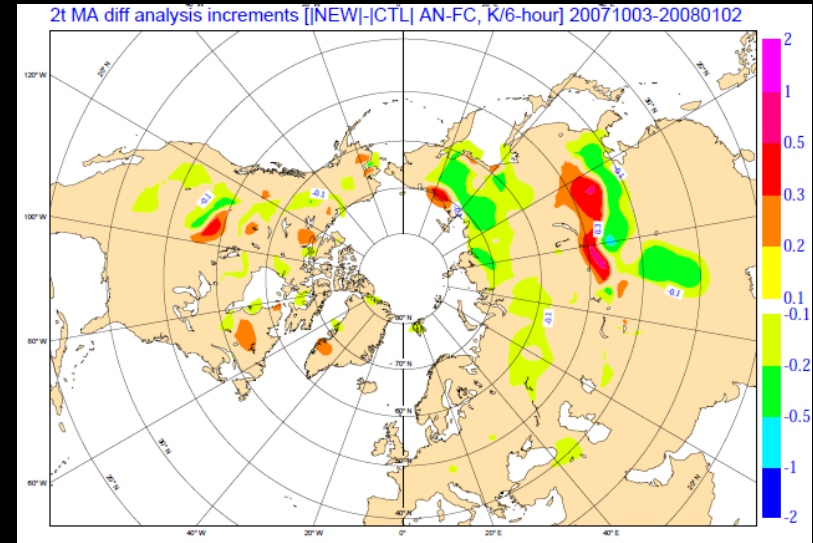
Global offline GSWP2 (1°x1°, 1986-1995)

# Extended data assimilation

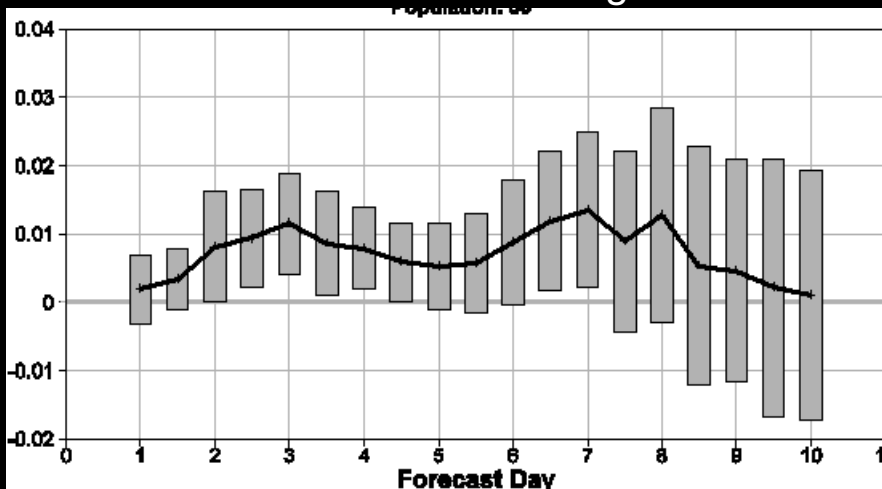
Snow mass analysis increments:  
|OPER| - |CTR|



2 m temperature analysis increments:  
|OPER| - |CTR|



Cold colors == reduction of assimilation increments ->  
Short range forecast closer to observations



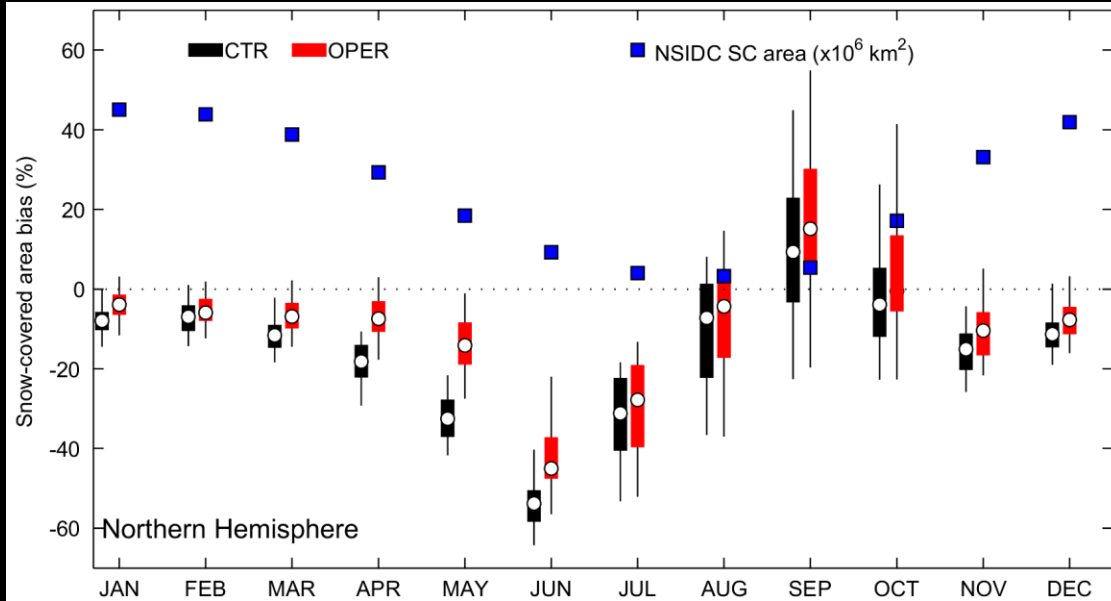
Root mean square error forecast (CTR-OPER)  
N. Hemisphere 1000hPa Temperature 00UTC

Significant improvement of near surface temperature  
up to day 7/8 of forecast

Set of 4d-var experiments :01-10-2007 -> 30-04-2008 (TL255L91)

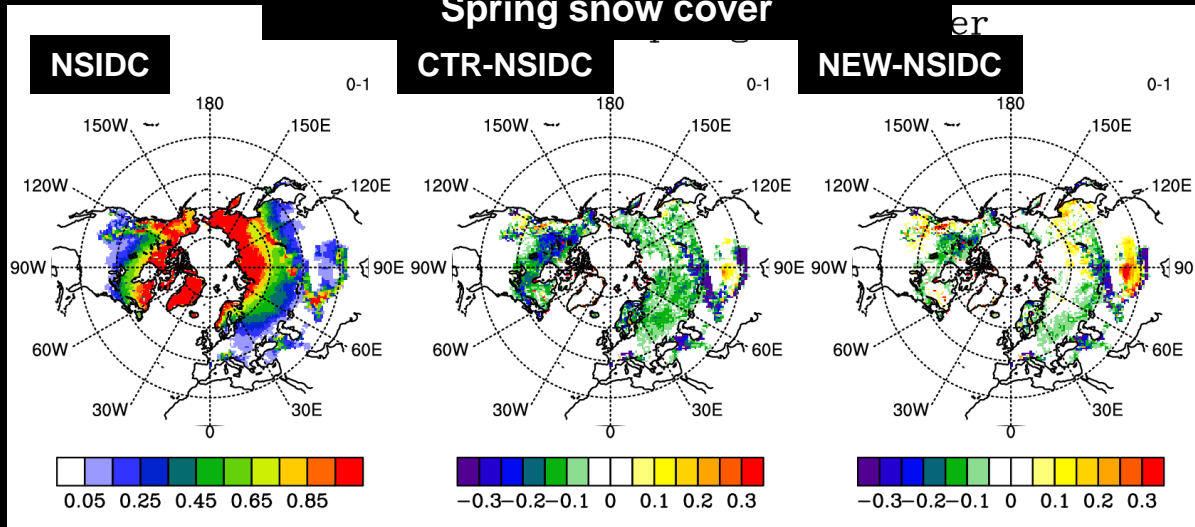
# Impact on EC-EARTH (snow cover)

## N. Hemisphere snow-covered area bias (% of snow-covered area NSIDC)



- CTR: under-estimation of snow-covered area from March- JUN
  - Early melting
- OPER:
  - Interception of rainfall on the snowpack
  - Revised snow cover fraction
  - Revised snow albedo
  - **Significant improvement of spring snow ablation**

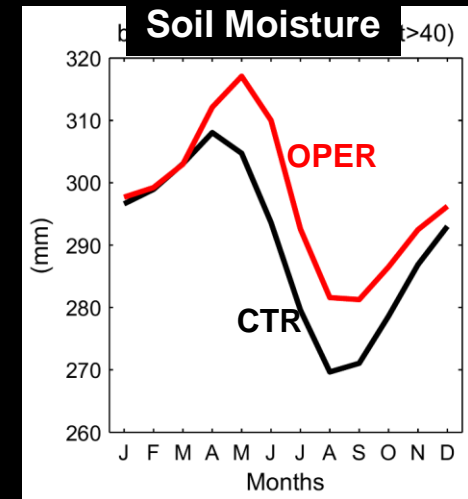
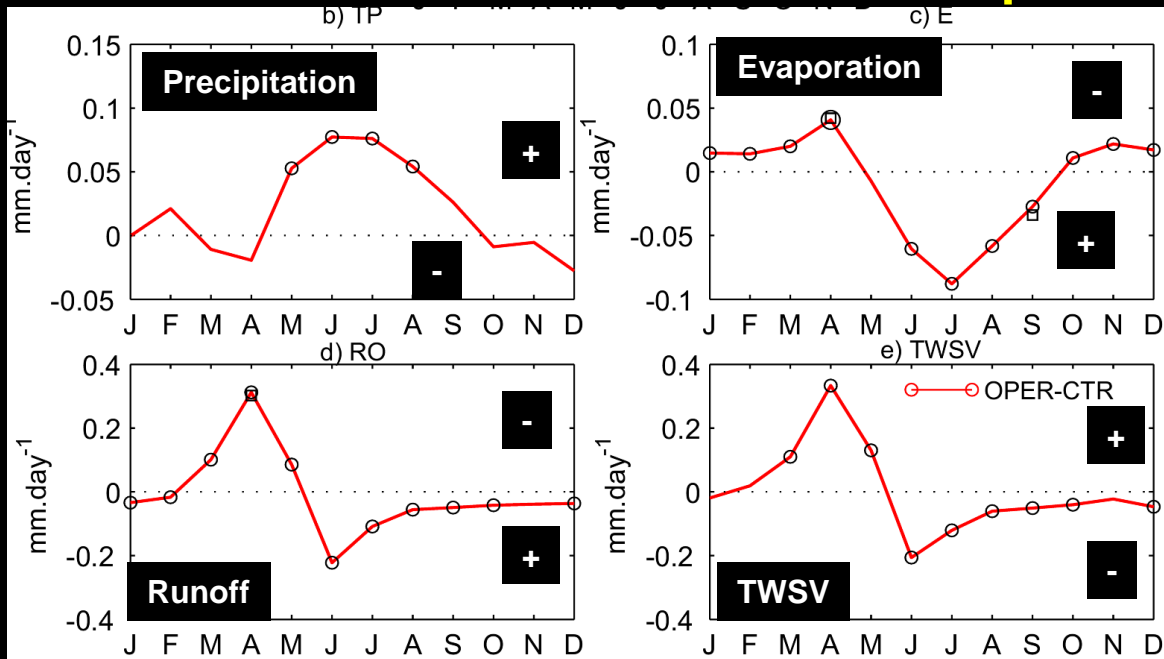
## Spring snow cover



Spring average Snow cover bias :  
 -CTR : -9.6 %  
 -OPER : -5.4 % (reduction of 43 %)



# Impact on EC-EARTH (water cycle)

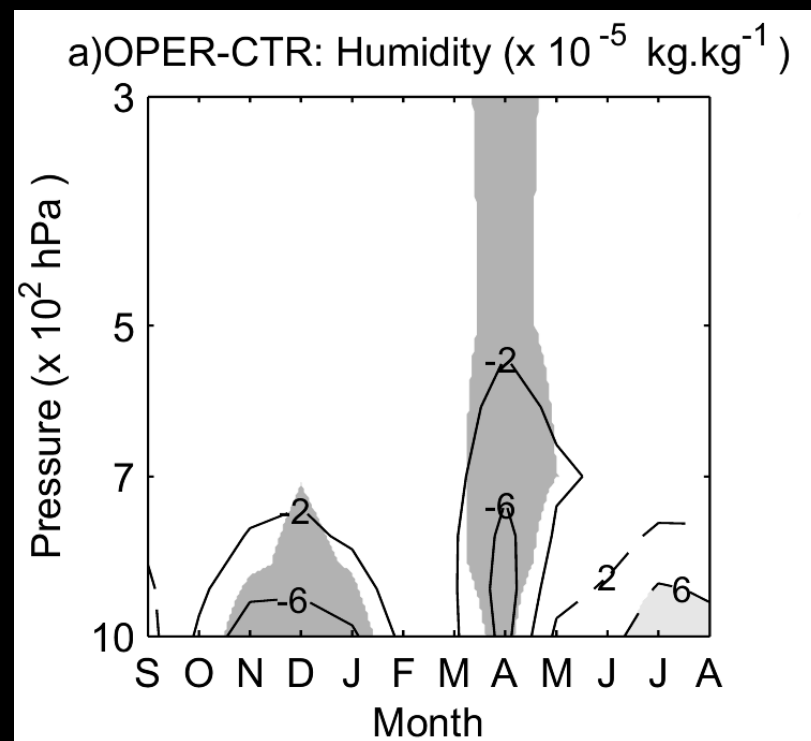
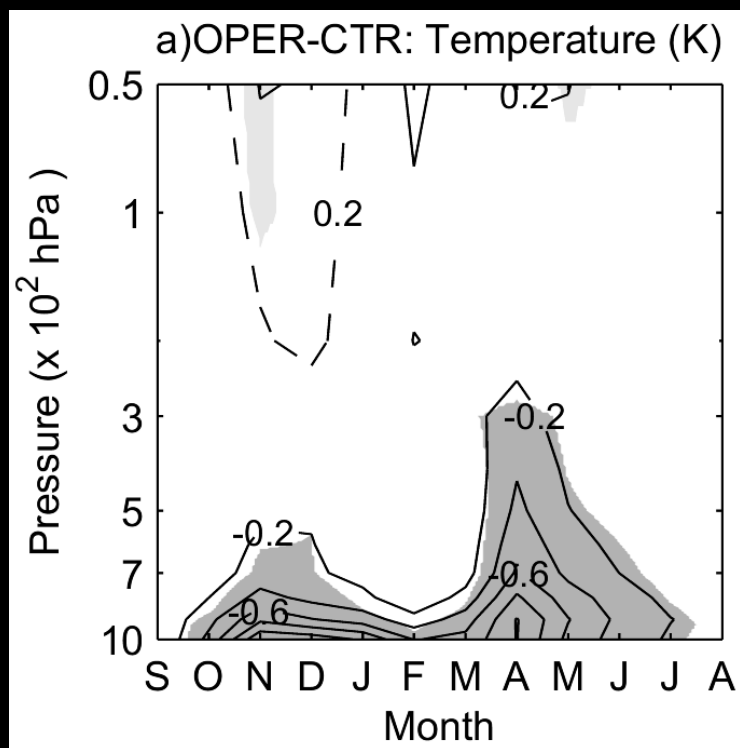


Averages over Eurasia land masses poleward 40° N.

Differences between OPER-CTR (symbols significant at 95 %)

- Increased soil moisture (Spring- Summer)
  - Reduction of early runoff (increased soil water storage)
  - + Evaporation during summer (+ soil moisture available)
  - + Precipitation during summer (more available humidity)
  - Intensification of the soil moisture / Precipitation feedback

## Impact on EC-EARTH (upper air)



Averages over N. Hemisphere polar cap lat  $>40^\circ$  N.

Differences between OPER-CTR (shaded: significant at 95% - dark gray OPER < CTR; light gray OPER > CTR)

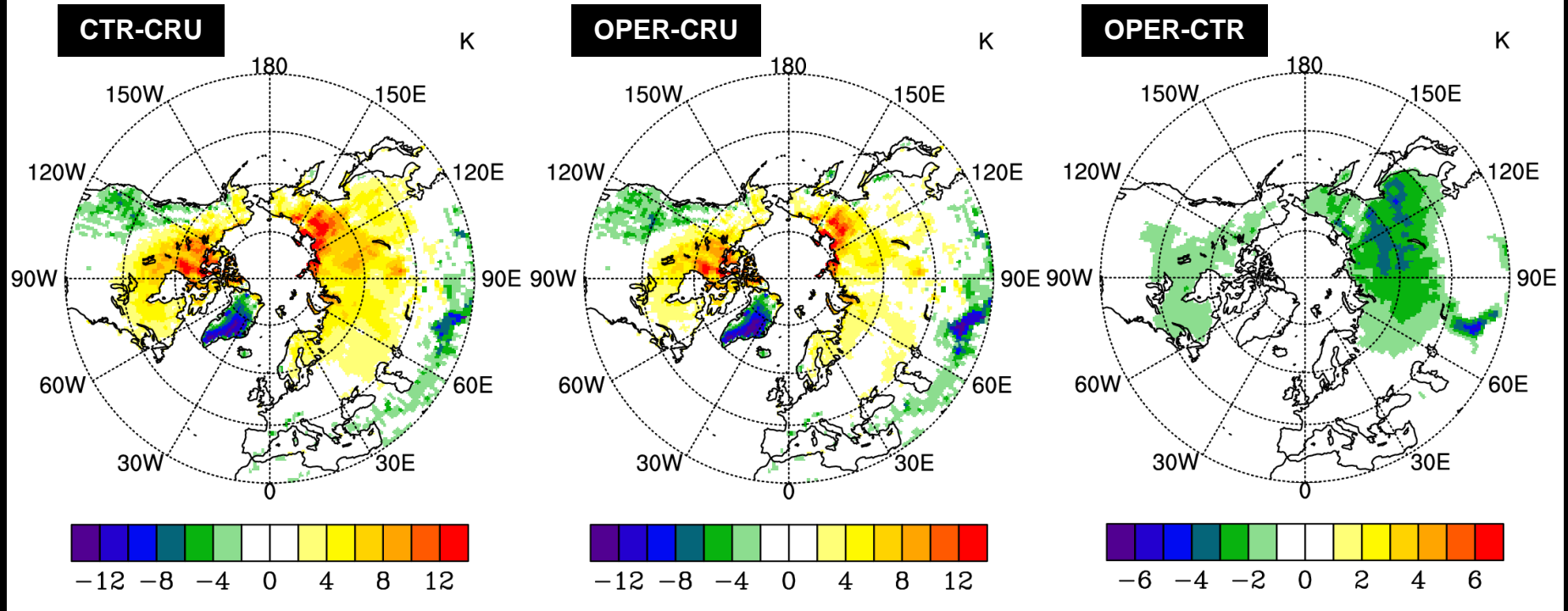
- Cooling of the troposphere up to 500 hPa during autumn/winter
- Cooling effect reaches 300 hPa during Spring
- Some warming in the top of the troposphere / stratosphere (not significant)

- Cooling of the troposphere + reduction of evaporation  $\rightarrow$  reduction of specific humidity
- Increased evaporation in summer  $\rightarrow$  increased humidity (restricted to the lower troposphere)

3x 30 yrs (1979-2008) EC-EARTH (atmosphere only with prescribed SSTs and sea ice)

# Impact on EC-EARTH (Temperature)

## Winter 2 meter temperature (K)



**CTR** – Warm bias over snow-covered areas (reaching 12 K)

**OPER**– Significant reduction of the warm bias (cooling between 4 and 6 K)

Increased snow insulation in OPER

-> decoupling between the PBL and underlying surface.

Winter Mean Absolute Error (Eurasia land masses poleward 60°N):

**CTR : 7.23 K**

**OPER : 6.29 K** (reduction of 13%)

**3x 30 yrs (1979-2008) EC-EARTH (atmosphere only with prescribed SSTs and sea ice)**

- Revised OPER snow scheme improved **local site simulations** ( late/early snow melt in open/forest sites);
- **Increased snow insulation** (due to reduced snow density) improved significantly the **runoff on large scale basins**;
- 4d-var assimilation tests showed a positive impact in both the **assimilation** and short range **weather forecast**;
- On climate EC-EARTH runs OPER:
  - Reduction of the **early snow melting** in the N. Hemisphere;
  - Intensification **soil moisture/precipitation** feedback (increased soil water storage in spring – summer);
  - Cooling of air temperature up to **500 hPa** polar cap >40°N;
  - Reduction of the **warm bias** in 2-meter temperature over snow-covered regions (stronger decoupling between the PBL and underlying soil)
- The OPER snow scheme was introduced in the ECMWF operational weather forecast in **September 2009 (CY35R3)**.

# Thank you

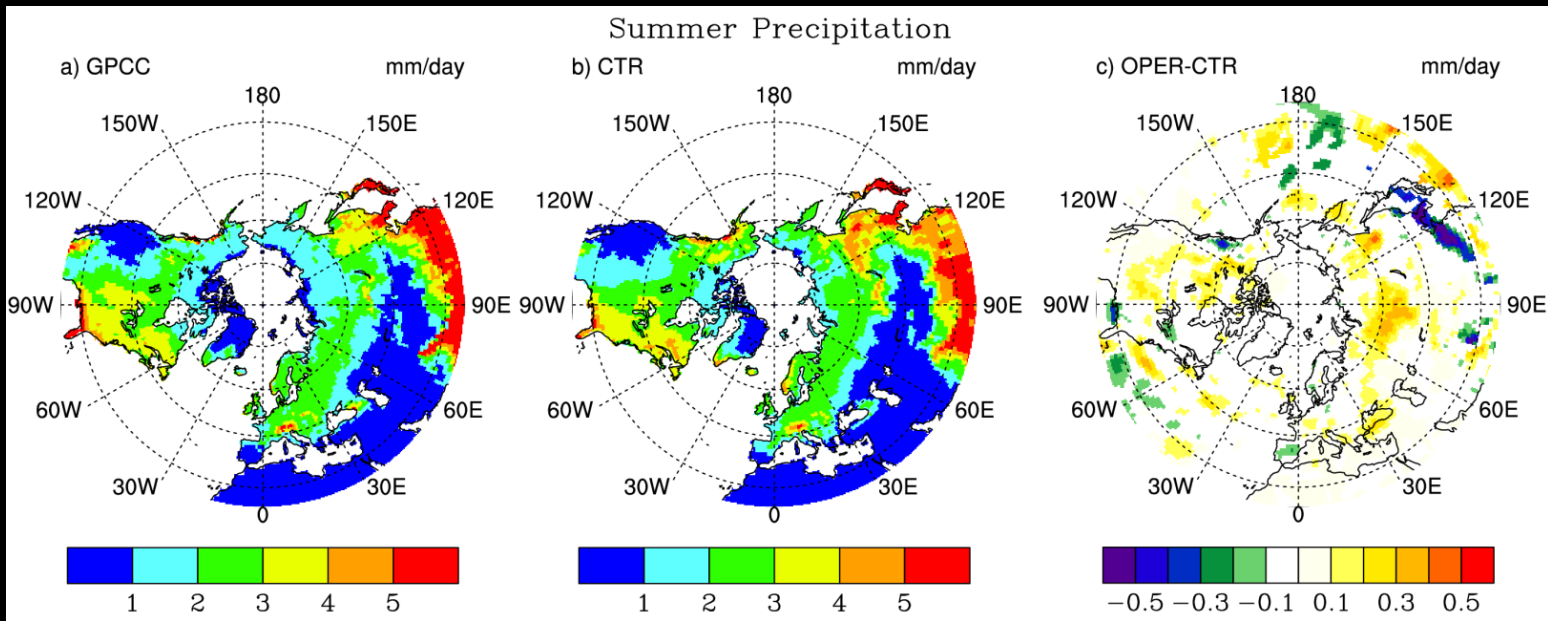
Dutra, E., Balsamo, G., Viterbo, P., Miranda, P. M. A., Beljaars, A., Schär C., and Elder, K., 2010: An improved snow scheme for the ECMWF land surface model: description and offline validation. **J. Hydrometeorol.**, doi:10.1175/2010JHM1249.1 (in press)

Also available as **ECMWF tech memo 607**

(<http://www.ecmwf.int/publications/library/do/references/show?id=89648>)

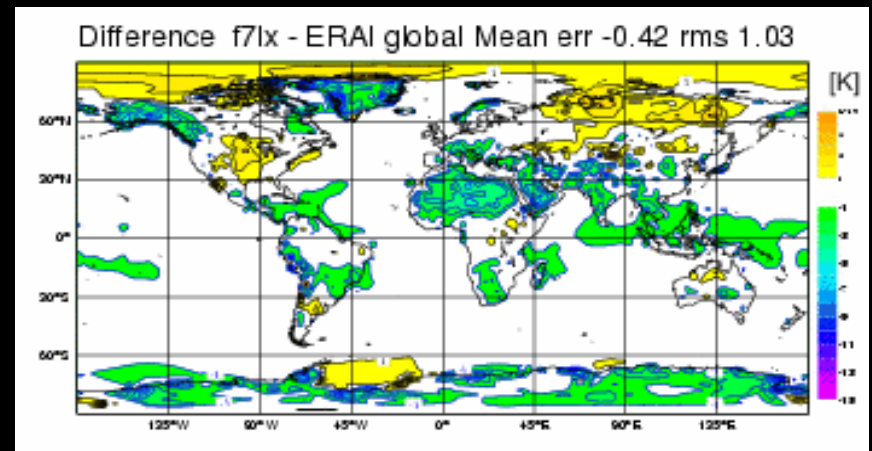
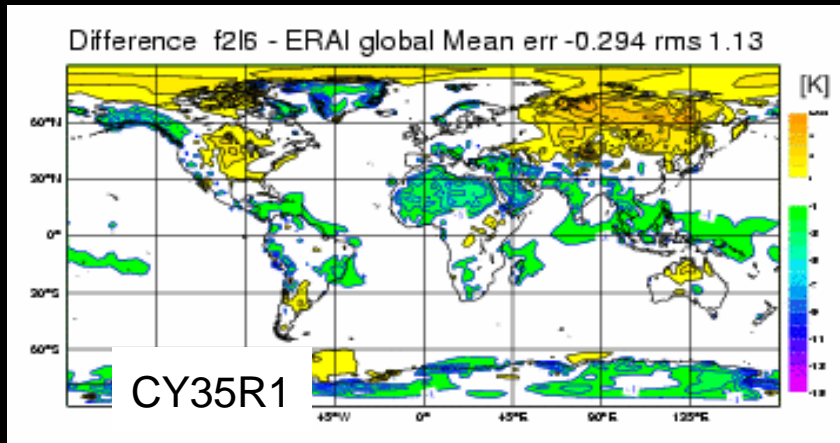
This work was supported by the Portuguese Foundation for Science and Technology (FCT) under project AMIC PTDC/AAC-CLI/109030/2008 cofinanced by the European Union under program FEDER. E. Dutra acknowledges the financial support of FCT under grant SFRH/BD/35789/2007 and Fundação Calouste Gulbenkian.





# Hindcasts (1 year climate runs)

The annual mean T2m bias (13-month 4-member hindcasts) is reduced in snow-areas





# New snow scheme: description

**CTR**

## Snow energy balance

**NEW** Dutra et al. 2010 J.H

Dry Snow

$$\left[ (\rho C)_{sn} D_{sn} + L_f S_i^c \frac{\partial f(T_{sn})}{\partial T_{sn}} \right] \frac{\partial T_{sn}}{\partial t} = R_{sn}^N - L_s E_{sn} - H_{sn} - G_{sn}^B - L_f M_{sn}$$

**Diagnostic of liquid water content** from snow mass and temperature:

- Additional snow heat capacity / **heat capacity barrier**

$$(\rho C)_{sn} D_{sn} \frac{\partial T_{sn}}{\partial t} = R_{sn}^N - L_s E_{sn} - H_{sn} - G_{sn}^B - L_f M_{sn}$$

## Snow mass balance

$$\frac{\partial S}{\partial t} = F - c_{sn} E_{sn} - M_{sn}$$

Rainfall bypasses the snowpack

$$\frac{\partial S}{\partial t} = F + c_{sn} F_l - c_{sn} E_{sn} - R_{sn}$$

**Interception of rainfall** on the snowpack

Intercepted water can freeze (warm the snow)

Exponential increase (100-300 kg m<sup>-3</sup>)

$$\rho_{sn}^{t+1} = (\rho_{sn}^* - \rho_{sn_{max}}) \exp(-\tau_f \Delta t / \tau_1) + \rho_{sn_{max}}$$

Fresh snow (snowfall) density  
cte = 100 kg m<sup>-3</sup>

## Snow Density

$$\frac{1}{\rho_{sn}} \frac{\partial \rho_{sn}}{\partial t} = \frac{\sigma_{sn}}{\eta_{sn}(T_{sn}, \rho_{sn})} + \xi_{sn}(T_{sn}, \rho_{sn}) + \frac{\max(0, Q_{sn}^{INT})}{L_f(S - S_l)}$$

Physically based (Anderson, 1976): Overburden, thermal metamorphism, Melting

Fresh snow (snowfall) density

$$\rho_{new} = a_{sn} + b_{sn}(T_{air} - T_f) + c_{sn}(V_a)^{1/2}$$

Boone and Etchevers 2001, J.H

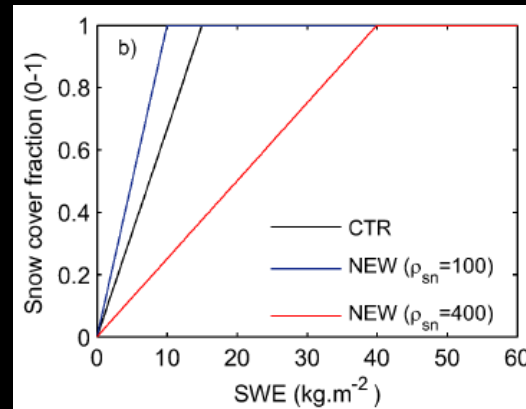
# New snow scheme: description

## CTR

$$c_{sn} = \min\left(1, \frac{S}{15}\right)$$

Only function of snow mass

## Snow cover



## NEW

$$c_{sn} = \min\left(1, \frac{S/\rho_{sn}}{0.1}\right)$$

Function of snow depth  
(snow mass and density)

## Snow Albedo

### Exposed areas

- Linear decay
- Exponential decay (Melting conditions)
- Reset to max (0.85) snowfall > 1 mm hr<sup>-1</sup>
- Linear decay
- Exponential decay (presence of liquid water)
- Continuous reset to maximum (0.85)

### Shaded snow (under high vegetation)

Constant = 0.15

- Vegetation type dependent based on MODIS data (Moody et al. 2007)

Index	Vegetation type	Albedo
3	Evergreen needle leaf trees	0.27
4	Deciduous needle leaf trees	0.33
5	Deciduous broad leaf trees	0.31
6	Evergreen broad leaf trees	0.38
18	Mixed forest / woodland	0.29
19	Interrupted forest	0.29