

# Retrieval of Snow Parameters from Ku- and X-Band Radar Backscatter Measurements

Helmut Rott<sup>1</sup>, D. Cline<sup>2</sup>, P. Etchevers<sup>3</sup>, I. Hajnsek<sup>4</sup>, M. Kern<sup>5</sup>, G. Macelloni<sup>6</sup>, E. Malnes<sup>7</sup>, J. Pulliainen<sup>8</sup>, S. Yueh<sup>9</sup>

*1 University of Innsbruck & ENVEO IT, Austria*

*2 NOAA-NOHRSC, Chanhassan, USA*

*3 Meteo-France, Saint Martin d'Hères, France*

*4 DLR-HR, Oberpfaffenhofen. Germany*

*5 ESA-ESTEC, Noordwijk, NL*

*6 IFAC-CNR, Firenze, Italy*

*7 NORUT IT, Tromsø, Norway*

*8 Finish Meteorological Institute, Helsinki, Finland*

*9 JPL-Caltech, Pasadena, USA*

# Outline of Presentation

---

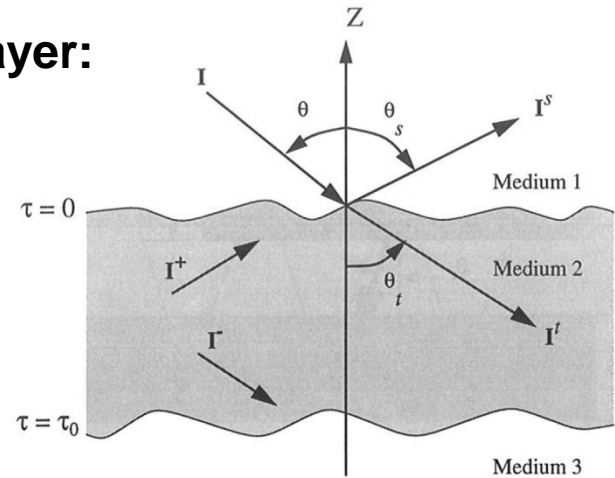
- Radiative transfer formulation for radar backscatter of snow-covered terrain
- Backscatter sensitivity of Ku- and X-band backscatter for SWE over ground
- The inversion problem for retrieving snow mass (SWE) of dry snow over ground and winter snow on glaciers
- Radiative transfer formulation and backscatter sensitivity for snow accumulation on glaciers
- Processing line for SWE retrieval from satellite-based SAR measurements
- Auxiliary data for segmentation and retrieval initialization
- Application example for retrieval algorithm
- Conclusions

# RT Transfer for SWE Retrieval

## Formulation for scattered intensity of single layer:

$$\mathbf{I}^S = \frac{1}{4\pi} \int_{4\pi} \mathbf{S}_T(\theta_s, \theta; \phi_s, -\phi) \mathbf{I}(\theta, \phi) d\Omega$$

$\mathbf{S}$  – Total scattering phase matrix of the layer accounting for scattering at interfaces, in volume, volume-interfaces, and multiple scattering in layer



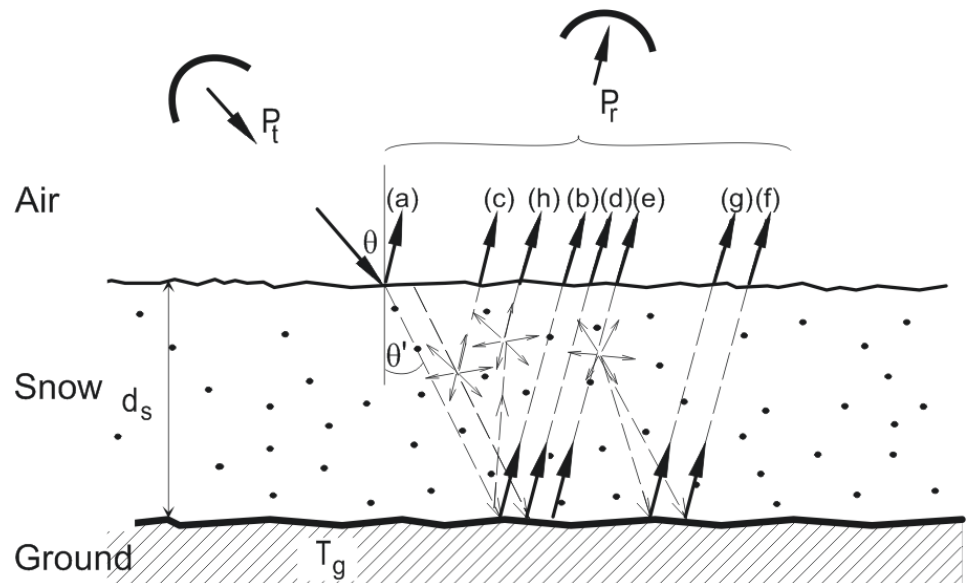
## Second order RT approach:

### Coherent backscatter terms:

- Direct at interfaces (a), (b)
- Direct volume (c), without multiple

### Incoherent backscatter terms:

- Multiple scattering in volume of (c)
- Volume/ground & ground/volume (ground surface separated in coherent and non-coherent terms)



# Semi-empirical RT-Formulation for SWE Retrieval

For iterative retrieval of physical parameters the number of free variables needs to be small, focussing at the main factors (multiple scattering lumped in  $\sigma^V$ )

The basic equation accounts for the *coherent* contributions – Single Layer:

$$\sigma_{f,pq}^t(\theta_i) = \sigma_{f,pq}^{as}(\theta_i) + \sigma_{f,pq}^V(\theta_t) + \left(T_f^{as}(\theta_t)\right)^2 \left[ \frac{\sigma_{f,pq}^G(\theta_t)}{L_{f,pq}^2(\theta_t)} \right]$$

One-Way Loss Factor:

$$L(\theta_t) = \exp(k_e d_s \sec \theta_t) = \exp(k'_e SWE \sec \theta_t)$$

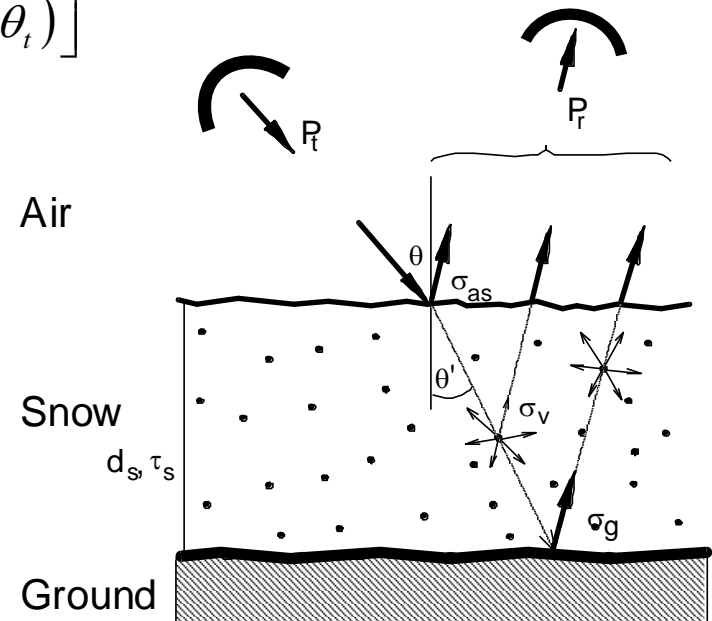
$$SWE = \rho_s d_s$$

$$\text{Extinction coefficient: } k_e = k_s + k_a \quad ;$$

$$\text{Density-normalized: } k'_e = k_e / \rho_s \quad k'_a = k_a / \rho_s$$

Volume scattering contribution:

$$\sigma_{f,pq}^V(\theta_t) = T_f^2(\theta_t) \left[ \frac{\omega_{f,pq}}{2} \left( 1 - \frac{1}{L_{f,pq}^2(\theta_t)} \right) \cos(\theta_t) \right]$$



Scattering

$$\text{Scattering albedo } \omega = k_s / (k_a + k_s)$$

# RT-Formulation for SWE Retrieval, Dual-Layer

Semi-empirical RT formulation for dual layer snow model

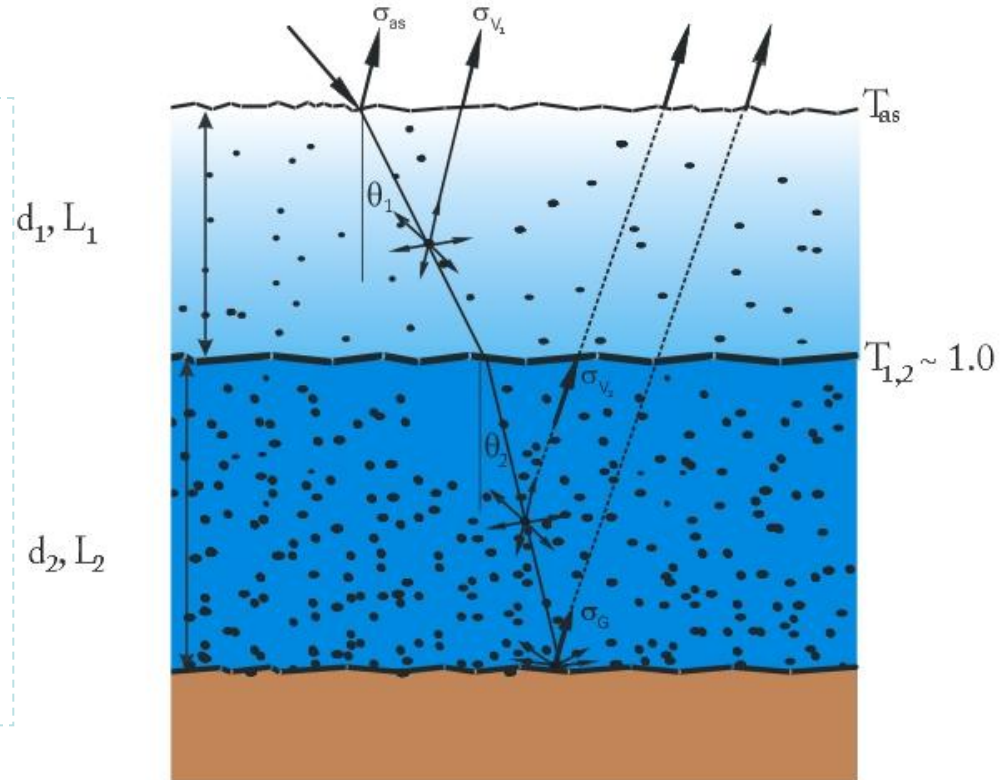
(Multiple interactions not specified, lumped in observed signals)

$$\sigma_{f,pq}^t(\theta_i) = \sigma_{f,pq}^{as}(\theta_i) + \sigma_{f,pq}^{V1}(\theta_{t1}) + \sigma_{f,pq}^{V2}(\theta_{t2}) L_1^2(\theta_1) + (T_f^{as}(\theta_{t1}))^2 \left[ \frac{\sigma_{f,pq}^G(\theta_t)}{L_1^2(\theta_1) L_2^2(\theta_2)} \right]$$

L – Loss factor

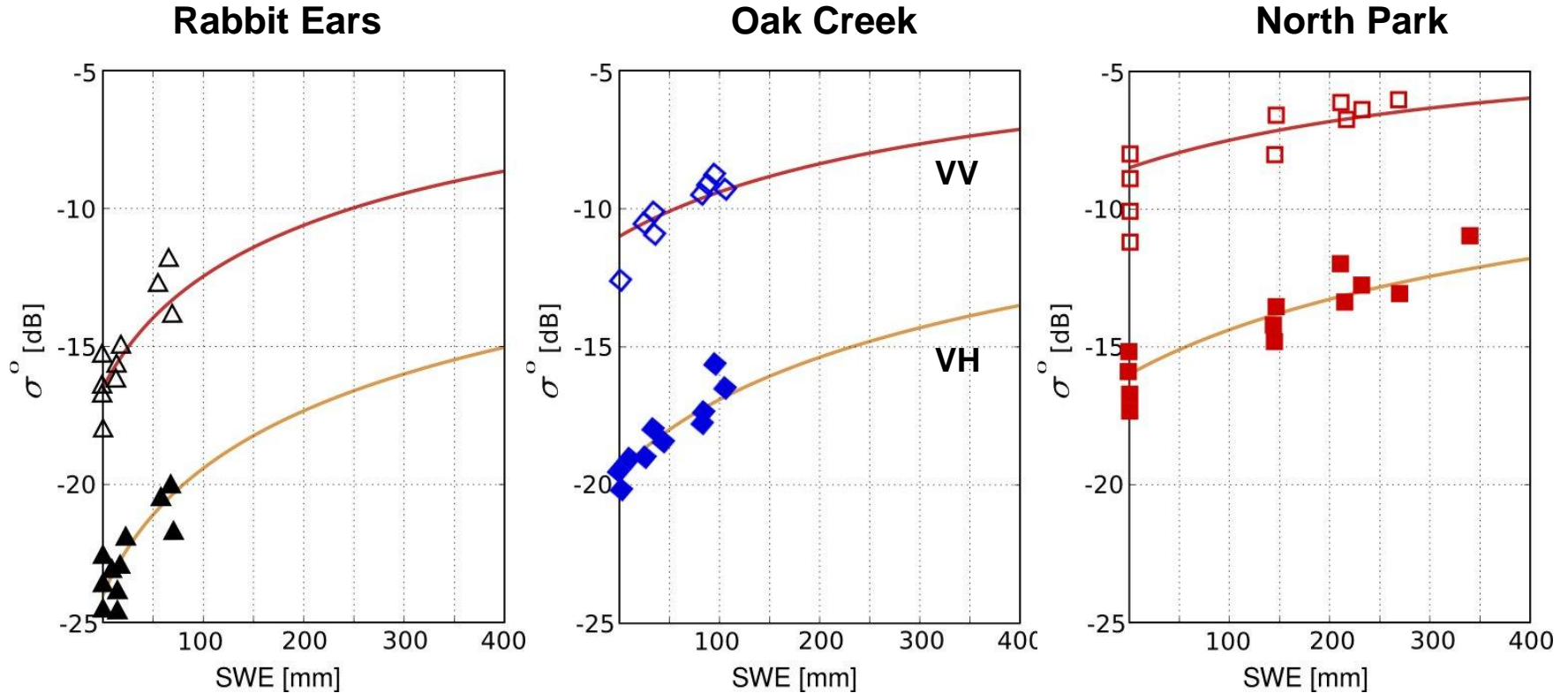
Strategy for inversion:

- Layer 1 represents the top snow layer governing the mass and energy fluxes with the atmosphere
- For cold snow Layer 2 properties change slowly by constructive metamorphism of grains → apply snow process model
- Iterate for SWE and  $\omega$  of Layer 1



# Backscatter Sensitivity to SWE – Variations of Volume Scattering Albedo and Background Signal

CLPX-II, Colorado, Winter 2006-07



PolScat  $\sigma^0$  (13.9 GHz) vs. in situ measurements of SWE

Forward computations with RT model,  $\omega = 0.80$ ,  $\omega = 0.75$ ,  $\omega = 0.70$

# Concept for SWE Retrieval

## Inversion of Physical Forward Model

Basic concept: Iterative matching of forward model and backscatter measurements to estimate free state variables

Input data requirements :

- Backscatter forward model
- Land cover data, DEM
- First guess snow parameters (from snow climatology or numerical meteorological data)

**Important:** Need for reliable and well tested forward model

**Advantages**

- Physical basis clearly defined
- Widely applicable

### Implementation options:

**Deterministic**

- Very sensitive to noise in measurements and model
- Solution may be non-unique

**Statistic**

- Unique solution, less sensitive to noise
- Reliable statistical data base on snow parameters needed

# Statistical Inversion for Retrieval of SWE

A semi-empirical radiative transfer model with reduced number of free parameters is used. The Nelder-Mead optimization algorithm is applied.

Cost Function: 
$$J(x) = \sum_{i=1}^n \frac{1}{2\sigma_i^2} [\Phi_i(x_1, \dots, x_q; c_{1i}, c_{2i}, \dots, c_{ri}) - Z_i]^2 + \sum_{j=1}^q \frac{1}{2\lambda_j^2} (x_j - x'_j)^2$$

Free parameters:

- SWE
- volume scattering coefficient  $k_s$  or scattering albedo  $\omega$

- $\Phi_i$  Forward RT model
- $Z_i$  Backscatter Measurement
- $\sigma$  Measurement noise
- $x_i$  State variables (1, ..., q)
- $c_i$  Configuration parameters
- $\lambda_i$  A priori standard deviation

## Variational Data Assimilation

$$J(\mathbf{x}) = (\mathbf{x} - \mathbf{x}_b)^T \mathbf{B}^{-1} (\mathbf{x} - \mathbf{x}_b) + (\mathbf{y}_{obs} - \mathbf{H}[\mathbf{x}])^T \mathbf{R}^{-1} (\mathbf{y}_{obs} - \mathbf{H}[\mathbf{x}]); \nabla J(\mathbf{x}) \Rightarrow \min$$

Requires background field  $\mathbf{x}_b$  and variance (physical snow parameters)  
 $\mathbf{H}$  – transformation operator from state vector  $\mathbf{x}$  to observation vector  $\mathbf{y}$



# Statistical Inversion for Retrieval of SWE

## RT model with reduced set of free parameters for iteration

$$\sigma_{pq}^t(\theta_i) = T_{pq}^2(\theta_t) \left[ \frac{\omega_{pq}}{2} \cos(\theta_t) \left\{ 1 - \exp\left(\frac{-2k'_e SWE}{\cos \theta_t}\right) \right\} + \sigma_{pq}^G(\theta_t) \exp\left(\frac{-2k'_e SWE}{\cos \theta_t}\right) \right]$$

$k'_e$  – density normalized extinction coefficient

Measurement vector of backscatter :

$$\sigma_i^t = \sum_{i=1}^4 \left[ \frac{\omega_i}{2} \cos \theta_t (1 - t_i^2) + \sigma_i^G t_i^2 \right]$$

Transmission coefficient

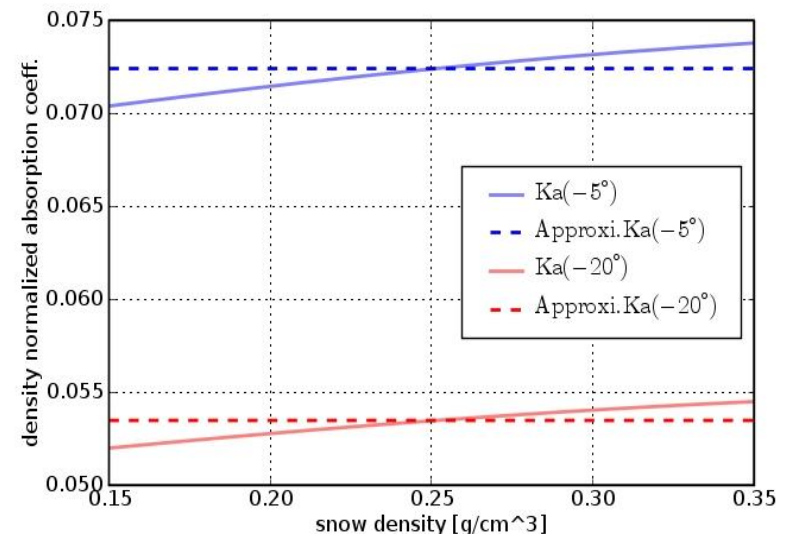
$$t_{i;i=1,2,3,4}^2 = \exp\left(\frac{-2(\omega_i + k'_{a,i=1,3})SWE}{\cos \theta_t}\right)$$

$k'_a$  – density normalized absorption coeff.

First guess values (time variable):

SWE (or depth),  $k_s(f_1)$ ,  $\langle T \rangle$

$\rho$  (for estimating dense medium effects)

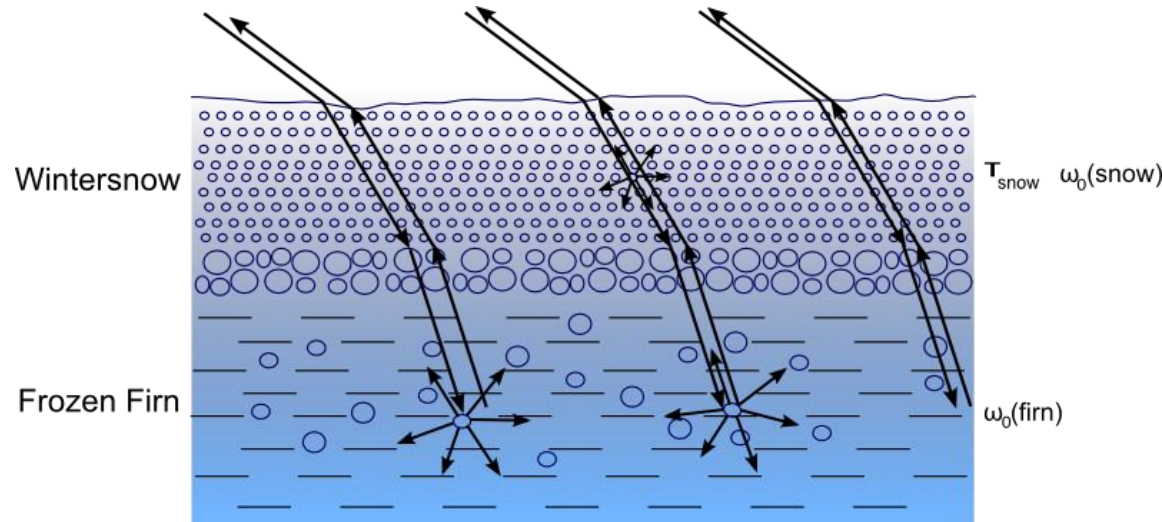


# RT-Formulation for Retrieval of Snow Accumulation on Glaciers

$$\sigma^t(\theta_i) = \sigma^{as}(\theta_i) + \sigma_{WS}^V(\theta_t) + T^2(\theta_t) \left\{ \sigma_{FI}^0(\theta_t) L_{WS}^2(\theta_t) \right\}$$

$\sigma_{WS}^V$  - volume scatter  
by winter snow

$$\omega_0(WS) < \omega_0(firn)$$

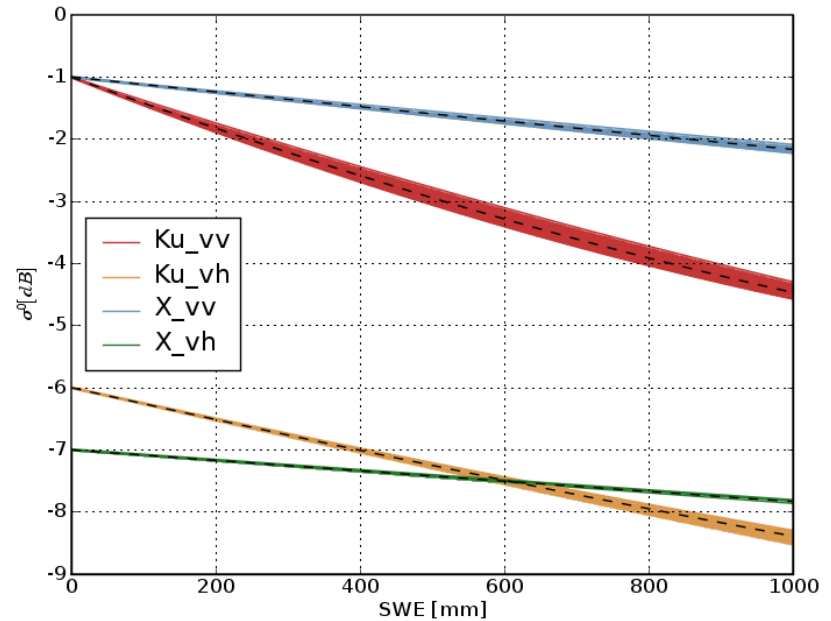
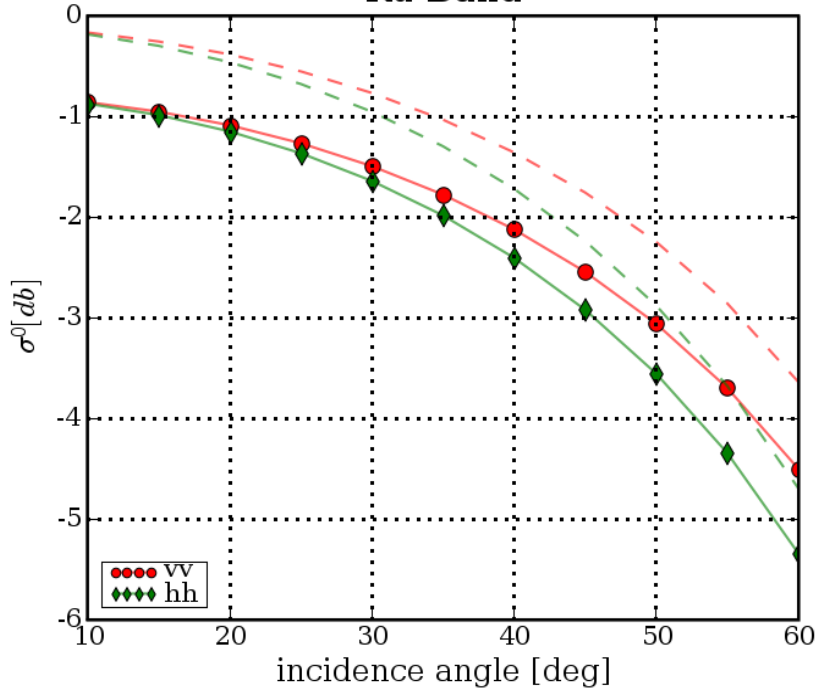


$$L_{WS}(\theta_t) = \exp(k_e d_{WS} \sec \theta_t) = \exp(k'_e SWE_{WS} \sec \theta_t)$$

$\sigma_{FI}$  - backscatter of frozen firn from early autumn period  
Iterate for  $SWE_{WS}$  by same procedure as for SWE on land

# Backscatter Sensitivity to Winter Snow Accumulation in the Percolation Zone of Glaciers

Ku-Band



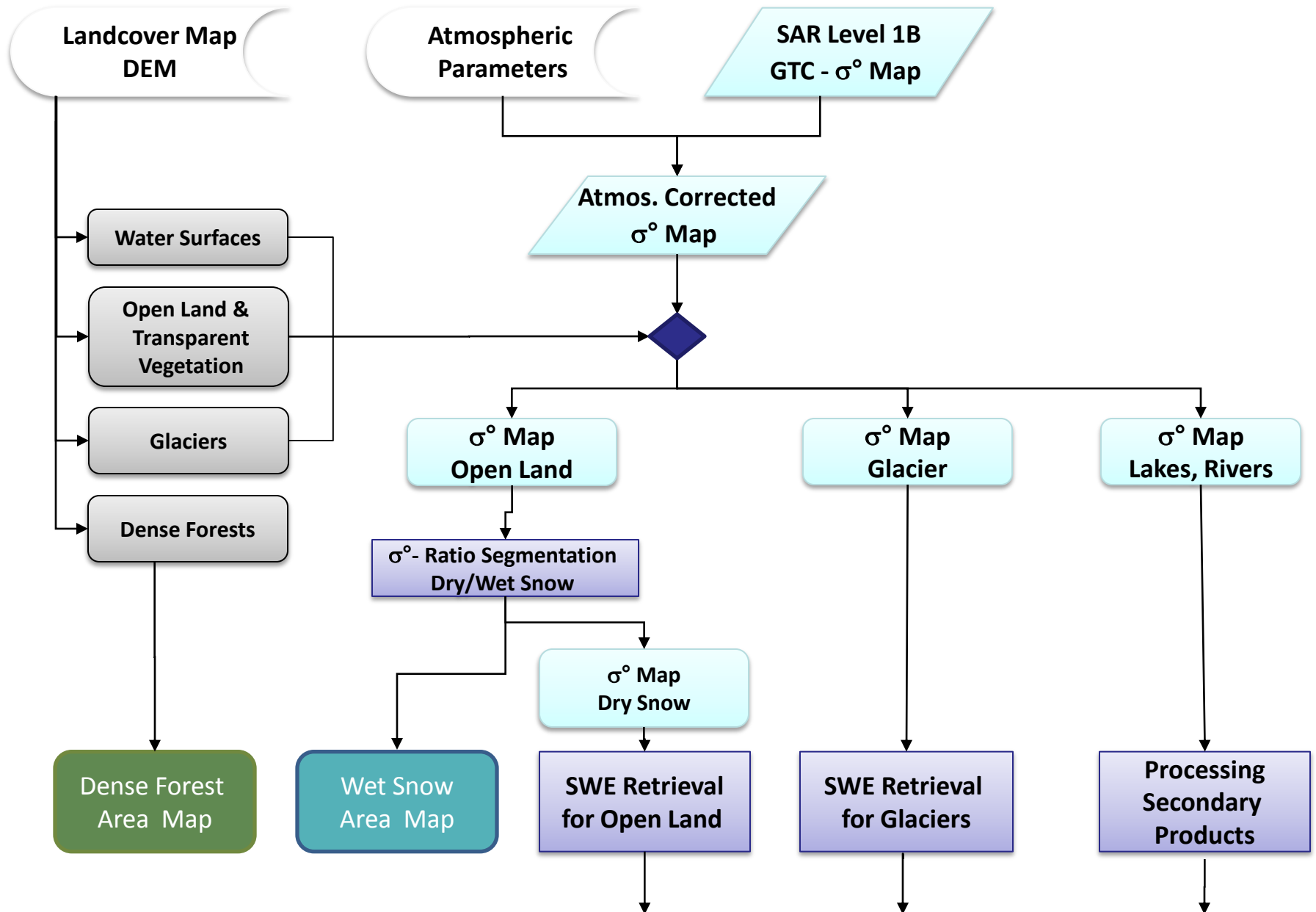
Broken Line – Frozen firn (early autumn)  
 Full Line – Firn and winter snow

Forward calculation of  $\sigma^0$  sensitivity to SWE, starting from  $\sigma^0$  observed in early autumn

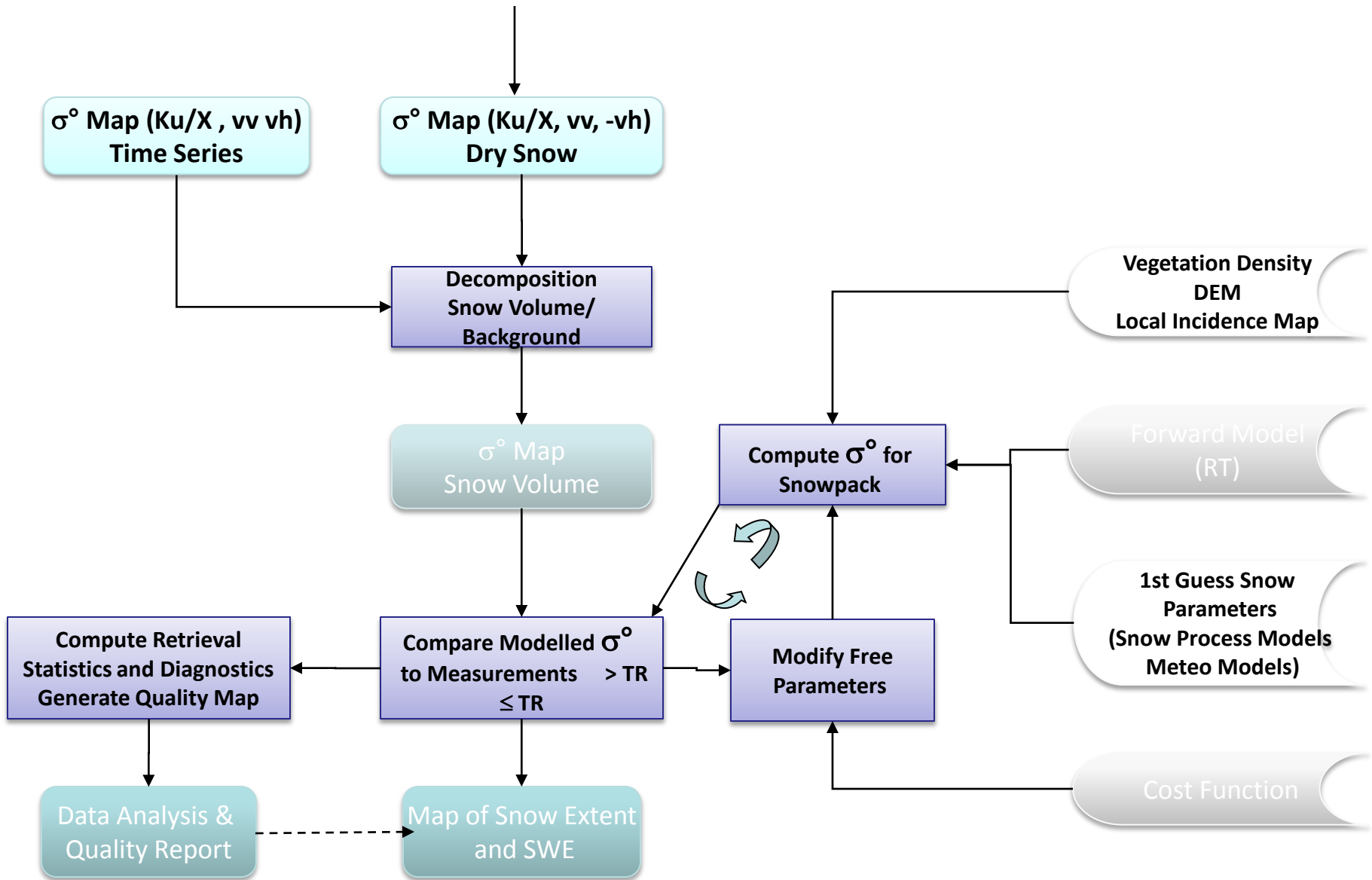
Winter snow

- Equiv. particle radius  $r = 0.2$  mm
- Ellipsoid axis ratio = 0.50
- Snow density =  $300 \text{ kg/m}^3$
- Snow depth = 1m

# SWE Processing Line – Part 1: Segmentation



# SWE Processing Line – Part 2: Iterative Retrieval

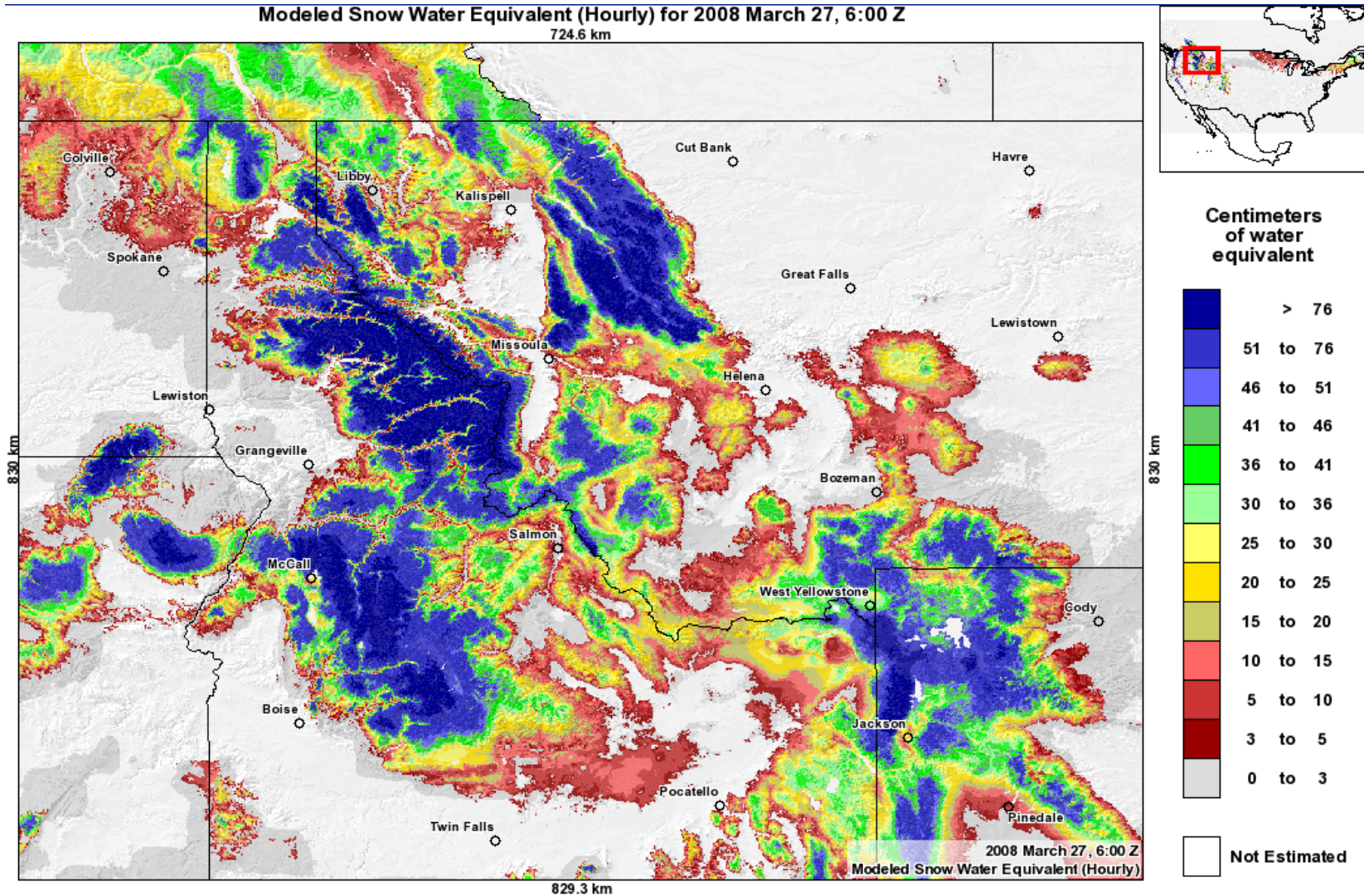


# Auxiliary Data for SWE Retrieval

- Atmospheric parameters for transmissivity correction (e.g. ECMWF analysis)
- Topography (DEM) – for geocoding and local incidence angle  
SRTM at  $\phi < 60^\circ$ ; ASTER GDEM; national ....
- Vegetation maps  
GlobCover (ESA, 300m), ECOCLIMAP, CORINNE (Europe), ...
- Background signature (snow-free case) from the SAR time series
- Snow statistics (climatology)
- 1<sup>st</sup> guess of snow parameters (e.g. gridded numerical meteo data from GCM and regional models, distributed snow process models, ...)
- *For glaciers*: Classification of diagenetic facies
  - Dry snow zone - Percolation zone - Glacier ice zone(Classification of facies by means of multi-temporal backscatter ratios)

# Option for 1<sup>st</sup> Guess Data - NOAA Snow Model

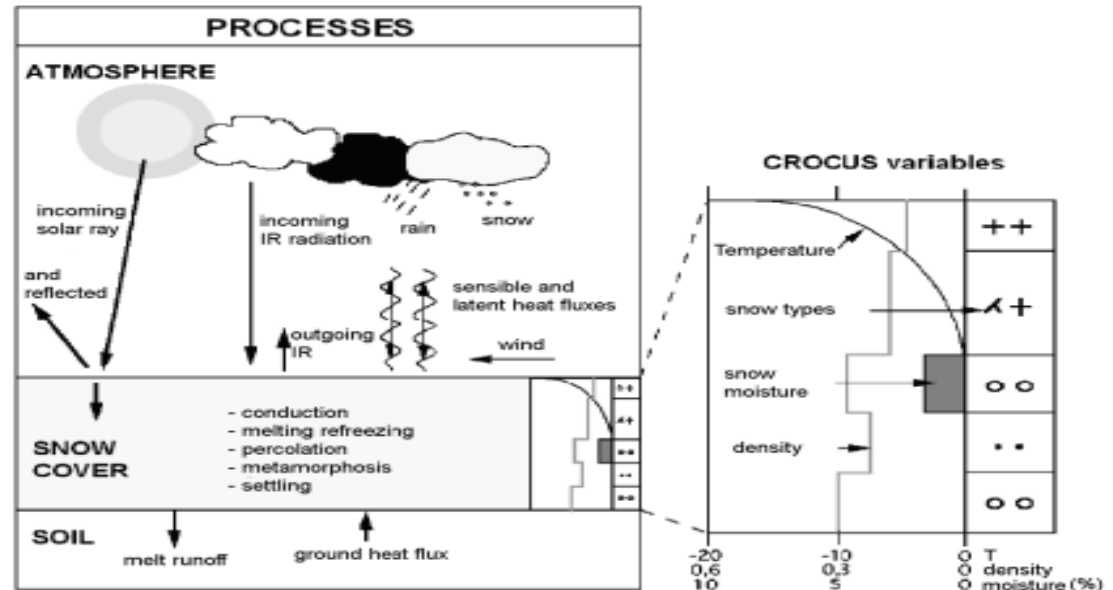
Maps on SWE, SD,  $\rho$ ,  $\langle T \rangle$  in 6 h time steps available



# Option for 1<sup>st</sup> Guess Data – SAFRAN/CROCUS

## SAFRAN/CROCUS

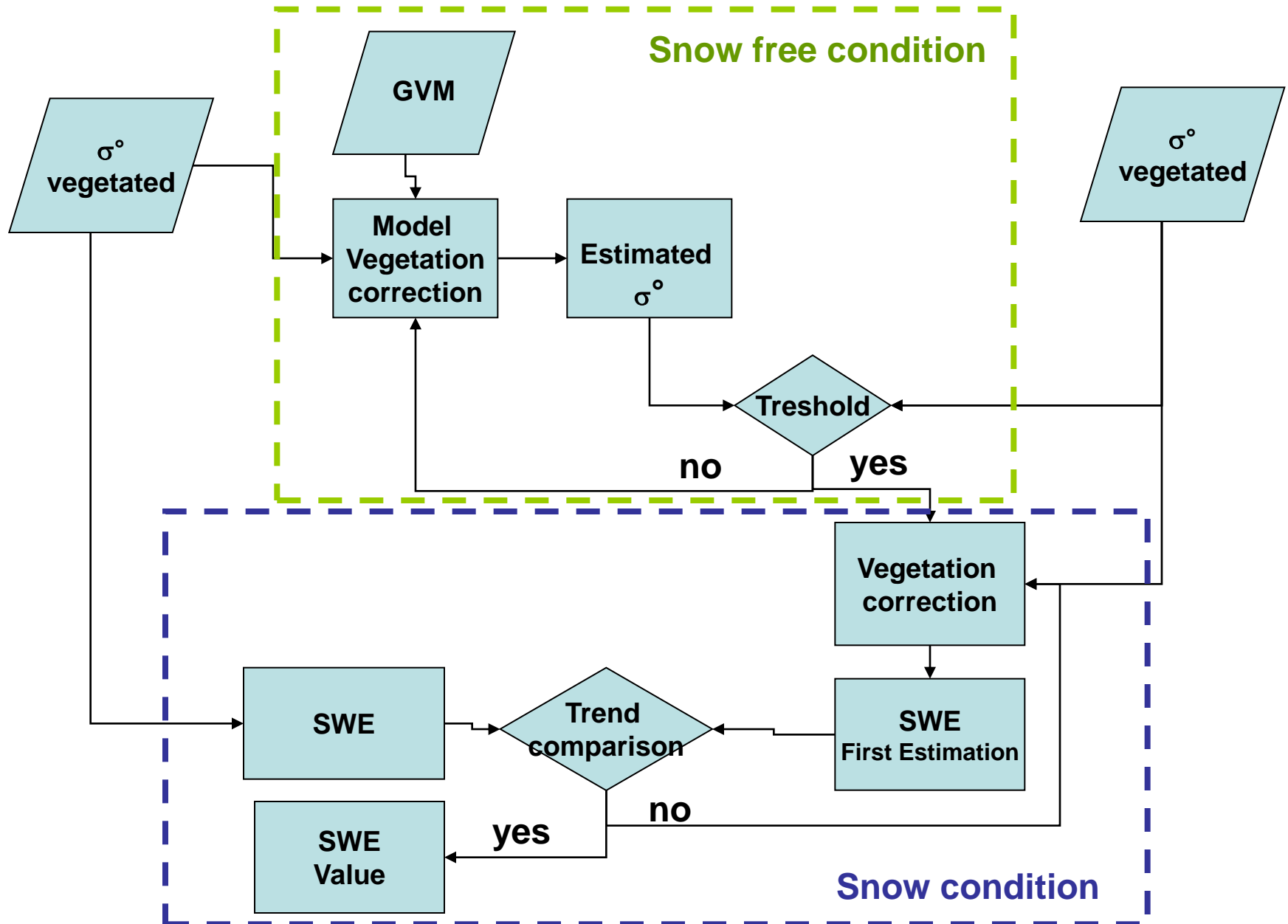
- SAFRAN delivers the fields of required meteorological input parameters (6 h steps)
- CROCUS computes the snow pack properties (mass, temperature, density, liquid water, layering)



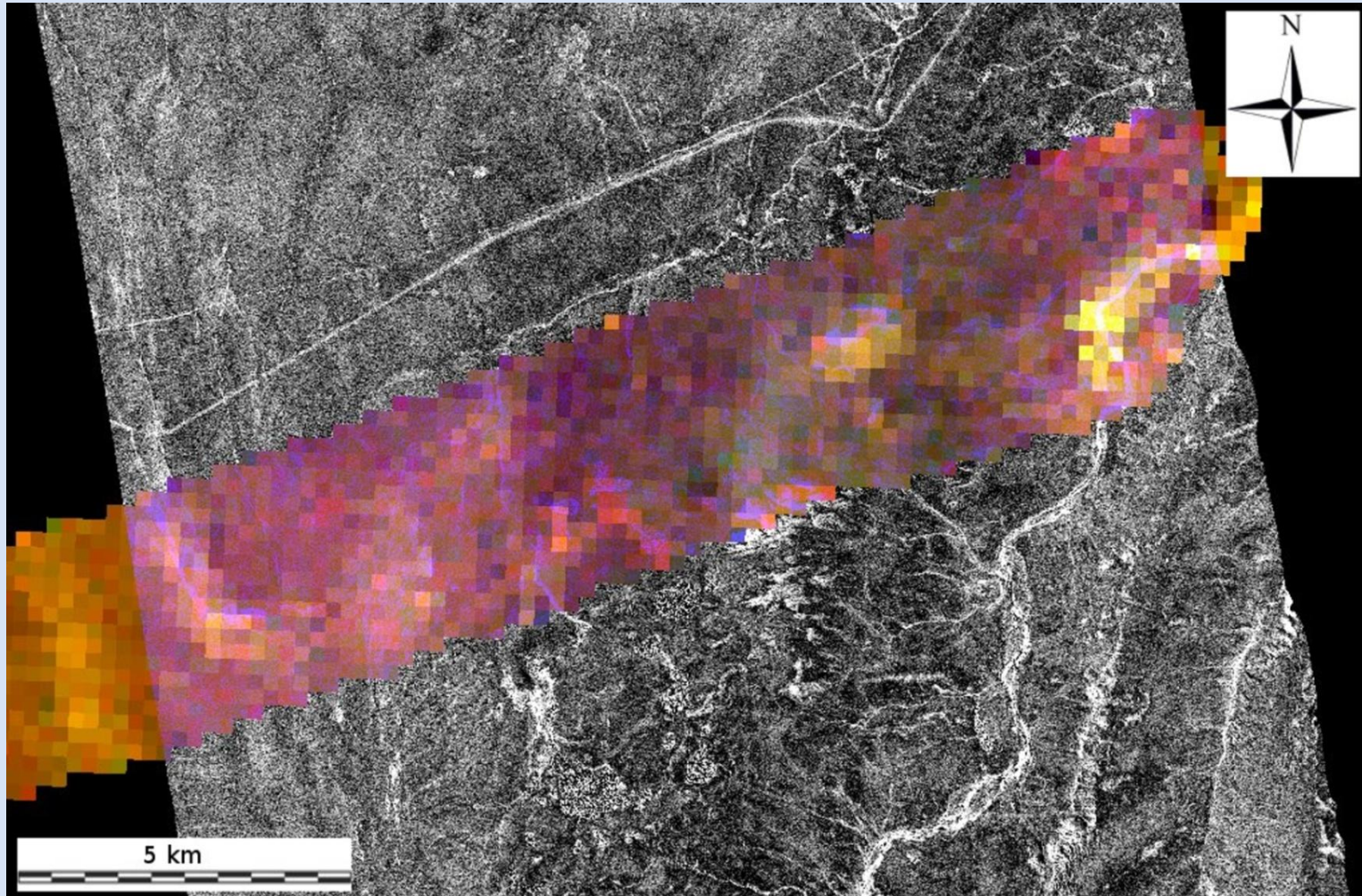
- Computes snowpack mass and properties in semi-distributed manner (elevation zones  $\nabla z = 300$  m , 6 orientations, 3 slope classes; local wind effects)
- Operationally used for snow hydrology, avalanche forecasting etc.



# Proposed Concept for Correction of Vegetation Effects



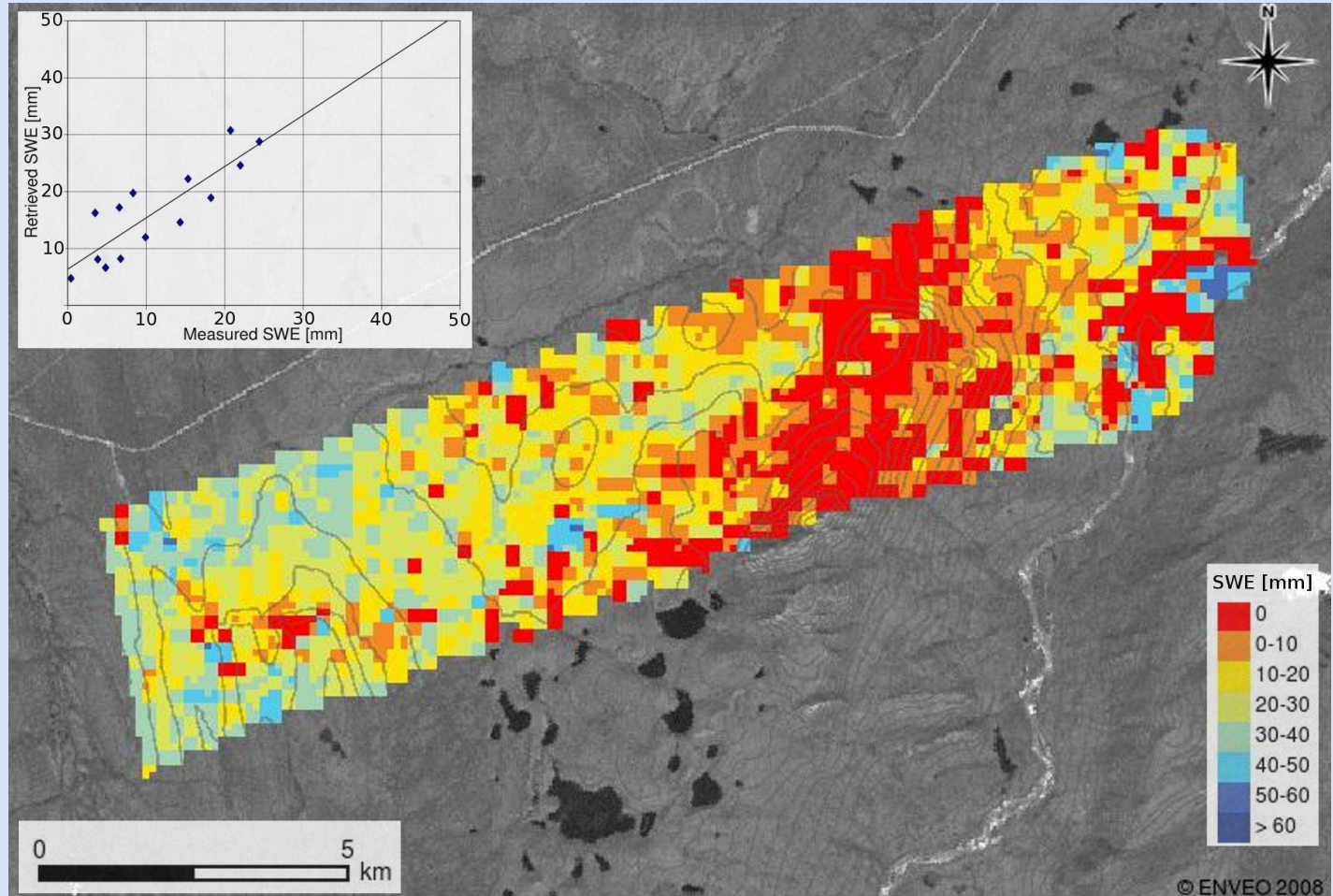
# Test of SWE Retrieval Concept



*CLPX-II Kuparuk River Study Area, Alaska, February 2008.  
Background: TerraSAR-X VV. Colours: Swath of PolScat on TerraSAR-X;  
Blue: X-band VV. Red: Ku-band VV. Green: Ku-band VH.*

# Test of SWE Retrieval Concept

Inset:  
Comparison of  
retrieved SWE  
vs. in-situ data



**SWE map** for Kuparuk River, Alaska, difference Feb. 2008 –Nov. 2007  
from Ku-band VV & VH (PolScat), X-band VV & VH (TerraSAR-X)

# Conclusion and Outlook

- Theory and experimental studies confirm the feasibility to apply dual frequency Ku-and X-band SAR measurements for mapping SWE
- Promising candidates for retrieval algorithms are statistical inversion techniques and variational data assimilation (VAR). They are robust and provide unique solutions
- The operator (based on RT model) linking observations (backscatter) to physical snow properties needs to avoid complexity, by still maintaining a realistic representation of the key processes
- Auxiliary data are important for
  - Segmentation (static data)
  - Corrections of vegetation effects, or masking of dense vegetation
  - Initialization of iterative retrieval (time-varying)
- Accounting for spatial and temporal context in the retrieval procedure should help to reduce noise
- For error assessment and validation of retrieval algorithms the acquisition of Ku-and X-band backscatter data over different snow regimes is of high priority